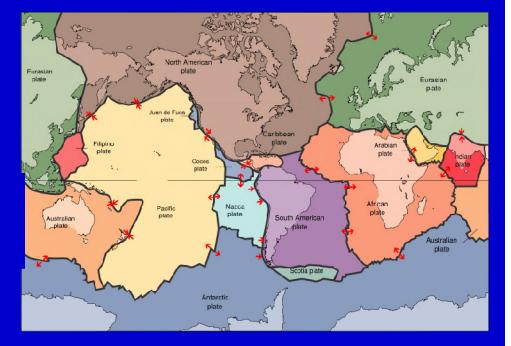
PLATE TECTONICS - Part I

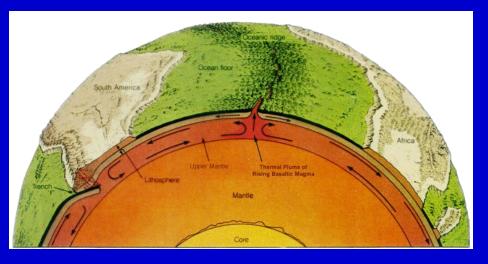
Geology's Modern Paradigm

Background and Overview



GEOL100 - Physical Geology Ray Rector - Instructor

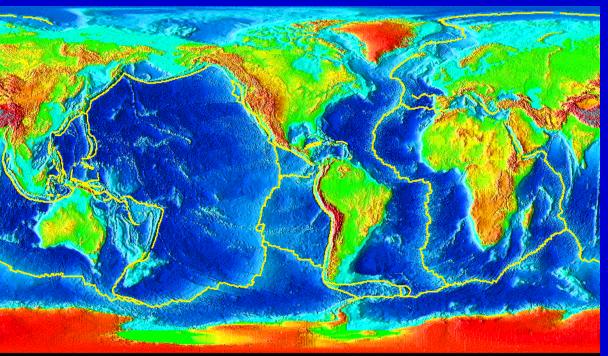




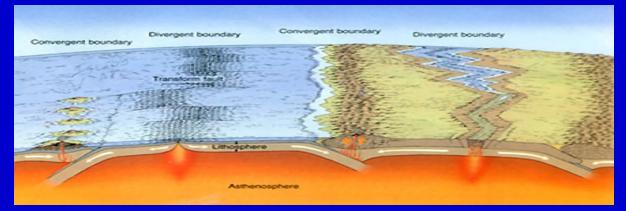
Topics in Plate Tectonics

<u>Topics</u>

- ✓ Age of the Earth
- ✓ Earth Physiology
- ✓ Isostasy
- ✓ Continental Drift
- ✓ Plate Tectonics Theory
- ✓ Seafloor Spreading
- ✓ Subduction
- ✓ Driving Mechanisms



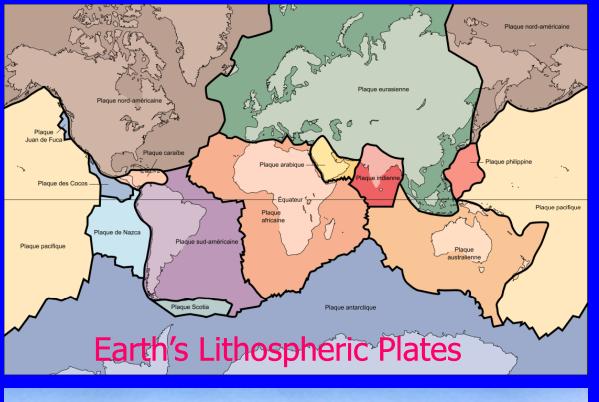
Crustal Plate Boundaries

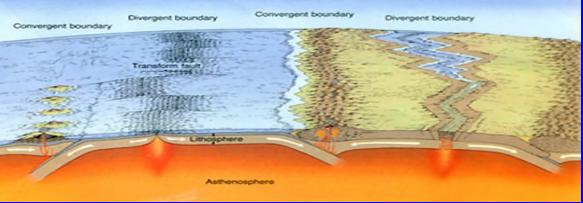


INTRO TO PLATE TECTONICS

Key Features:

- ✓ 14 Lithosphere Plates
- 6 Major, 8 Minor
- ✓ 100-300 km thick
- Strong and rigid
- ✓ Plates float on partially molten asthenosphere
- ✓ Plates are mobile
- ✓ Cm's/yr motion rates
- ✓ Seafloor Spreading creates new oceanic plates
- ✓ Subduction destroys older oceanic plates





Dynamics of a Restless Planet

Earth Exhibits a Long History of Mountain Building Events

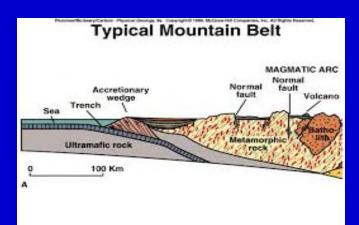
 ✓ Activity stretching over billions of years of time

 ✓ Numerous belt-like regions of exposed crustal rocks show intense deformation



Present-day Mountain Belt of Folded and Faulted Crust





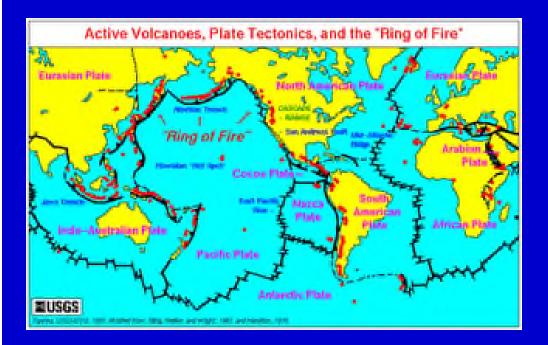
Dynamics of a Restless Planet

Earth's Surface Exhibits a Long History of Volcanic Activity

✓ Billions of years of volcanic activity

✓ Widespread evidence of regional-scale volcanism occur in belt-like exposures

✓ Volcanism found in both continental and oceanic settings





Mt St Helens Eruptions

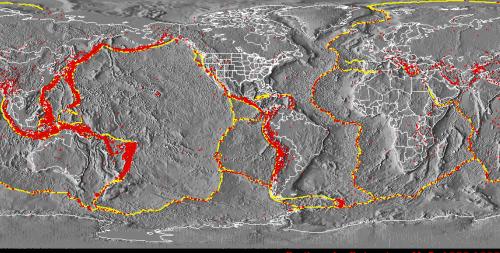
Dynamics of a Restless Planet

Earth's Surface Exhibits Extensive Faulting Activity

✓ Evidence of faulting stretching over billions of years of time

 ✓ Worldwide occurrence of local and regional-scale faulting occur along beltlike regions

✓ Faulting and associated quakes found in both continental and oceanic settings



Crustal Plate Boundaries Coastlines, Political Boundaries



The Great San Francisco Earthquake of 1906

How Old Is the Earth? How Can We Determine the Age of Earth? How Can We Date Earth's Geologic Events?

Earth's

Scientific Means of Dating Earth

Two Primary Means of Dating Rocks:

1) Relative Dating

- Determines the temporal order of rock forming events
- ✓ Does not give numeric ages
- ✓ Use of stratigraphic principles and fossils

2) Absolute Dating

- Determines the numeric age of rock forming events
- Only appropriate for ages of igneous rocks and minerals
- Primary method is the *radiometric technique*
- Used in conjunction with stratigraphic principles and fossils

Relative Dating and Stratigraphy

Relative Dating Principles

1) Superposition

- Oldest on bottom
- ✓ Youngest on top

2) Cross-cutting

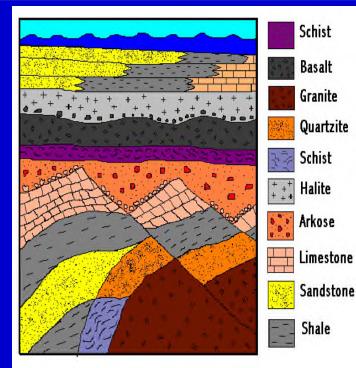
 Cross-cutting structure is younger than the structure that is being cross-cut

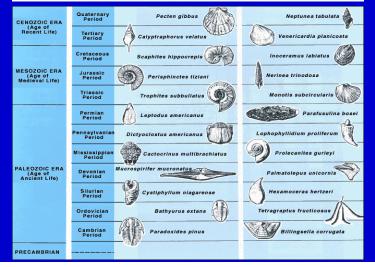
3) Inclusion

 Inclusion is older than rock that surrounds it

4) Fossil Succession

 Rocks containing a specific fossil species indicates a specific age





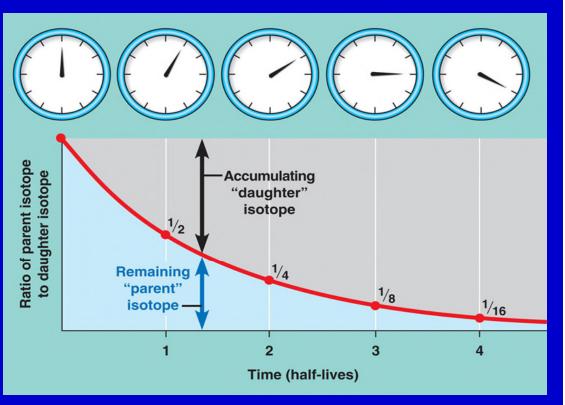
Principles of Radiometric Decay

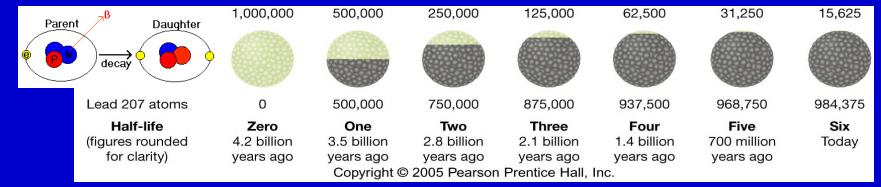
The Principles

 ✓ Spontaneous decay of unstable parent element into a its unique stable daughter element

✓ The half-life of each parentdaughter pair is a constant

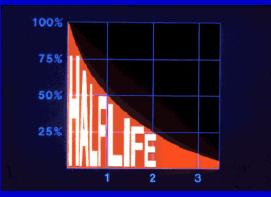
 ✓ Age of an igneous rock is determined by measuring the ratio of rock's parent-daughter material





Radiometric Half-Lives

Radioactive Parent/Daughter Pairs and Associated Half-Lives



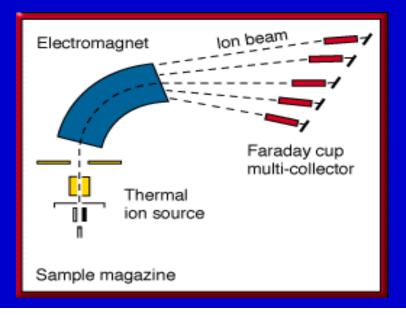
Parent Isotope	Stable Daughter Product	Currently Accepted Half-Life Values
Uranium-238	Lead-206	4.5 billion years
Uranium-235	Lead-207	704 million years
Thorium-232	Lead-208	14.0 billion years
Rubidium-87	Strontium-87	48.8 billion years
Potassium-40	Argon-40	1.25 billion years
Samarium-147	Neodymium-143	106 billion years

Radiometric Dating Method

Analysis of Parent/Daughter Isotopic Compositions in Rocks

- ✓ Parent and daughter elements are isolated and refined from host mineral using conventional wet chemistry methods.
- ✓ Geochronologists determine the isotopic abundances of each paired parent and daughter element using a mass spectrometer.
- ✓ Isotopic abundance data are then used to determine rock age using the decay formula.





Radiometric Dating Method

Radioactive Decay of Parent Isotope into a Daughter Isotope

The mathematical expression that relates radioactive decay to geologic time is called the *age equation*:

$$t = \frac{1}{\lambda} \ln \left(1 + \frac{D}{P} \right)$$

where t is the age of the rock or mineral specimen,

D is the number of atoms of a daughter product today,

P is the number of atoms of the parent isotope today,

In is the natural lograithm (logarithm to base e), and

λ is the appropriate decay constant.

(The decay constant for each parent isotope is related to its half-life,

$$t^{1/2}$$
 by the following expression: $t^{1/2} = \frac{\ln 2}{\lambda}$

Earth's Age - Radiometric Dating of Rocks

Earth's Oldest Rocks

106	Description	Technique	Age (in billions of years)
	Acosta Gneiss (NW Territories, Canada)	207Pb-206Pb isochron	4.031 ± 0.003

Oldest Moon Rocks

Mission	Technique	Age (in billions of years)
Apollo 17	Rb-Sr isochron	4.60 +- 0.1

Oldest Meteorites

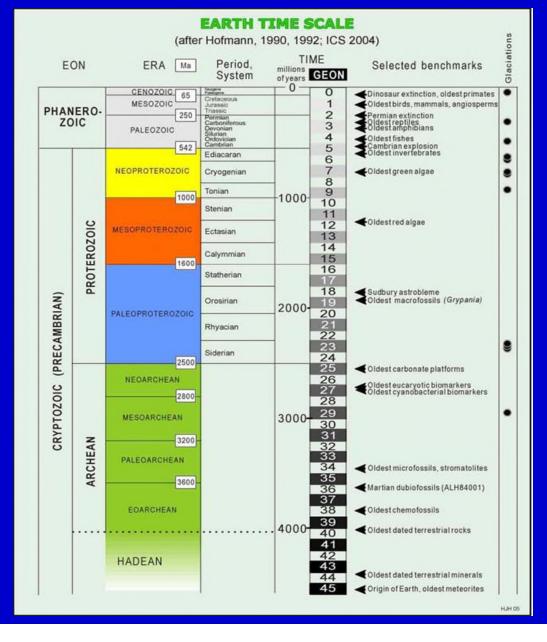


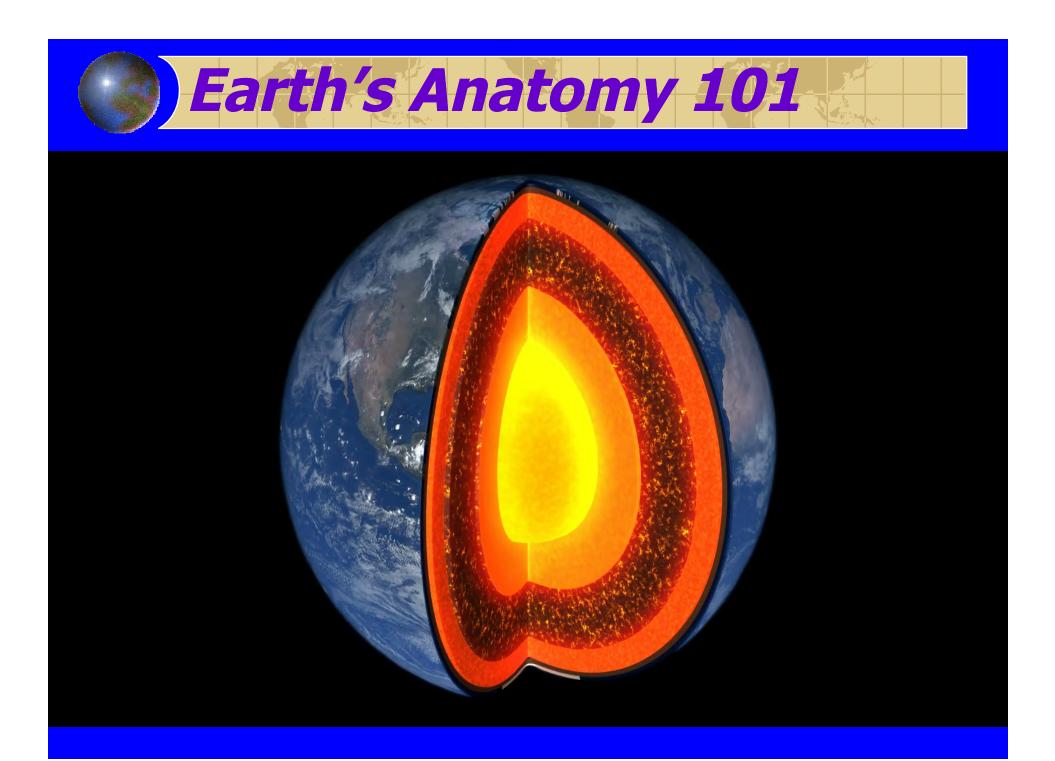
Description	Technique	Age (in billions of years)
Norton County (achondrite)	Mineral isochron	4.70 +- 0.1

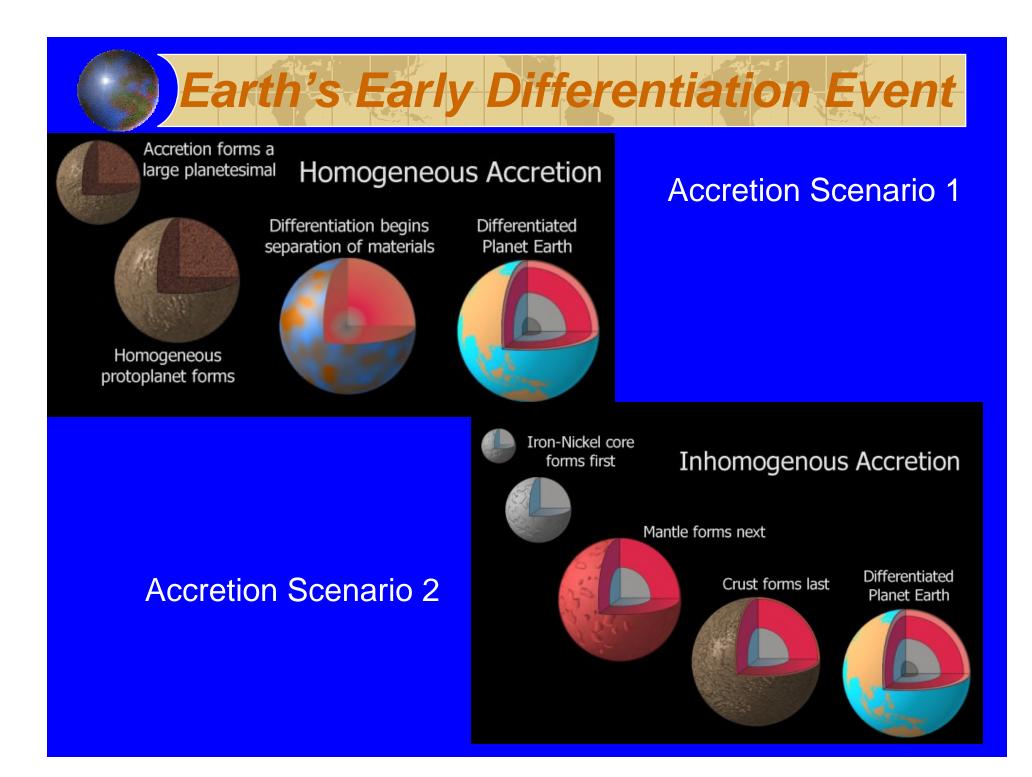
Earth's Geological Timescale

Key Ideas:

- 1) Originally based on relative dating and the use of age-specific fossils
- 2) Each period defined by unique assemblages of organisms
- Periods separated by mass extinction events
- 4) Numeric ages derived from radiometric analysis of igneous rocks found within the stratigraphic record

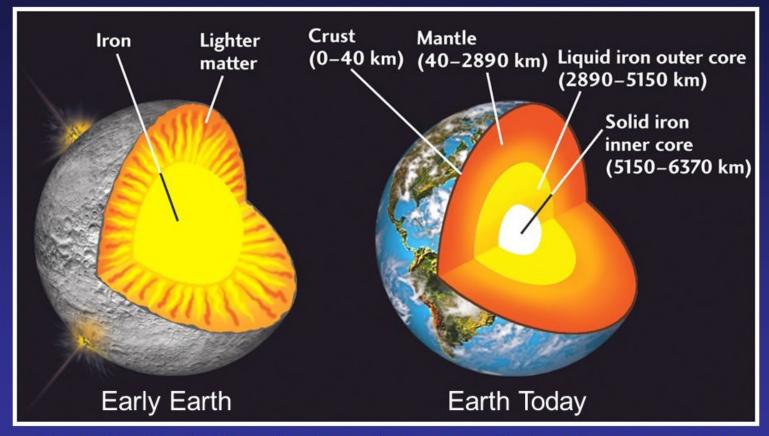




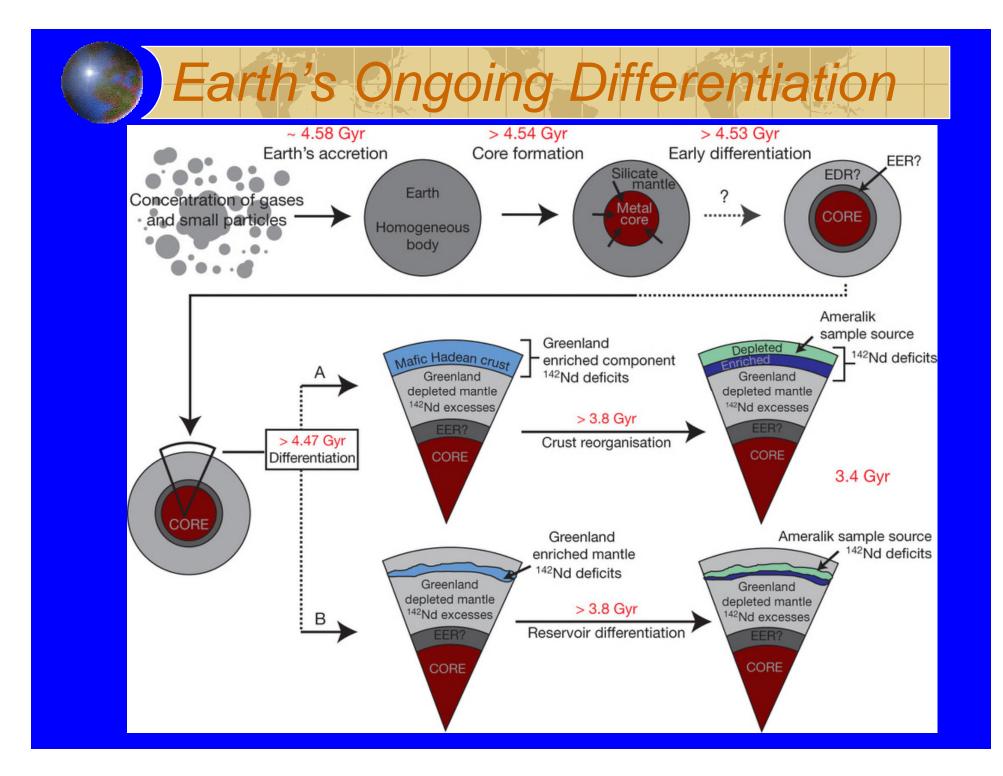


Chemical Differentiation of the Earth

Earth's History of Differentiation

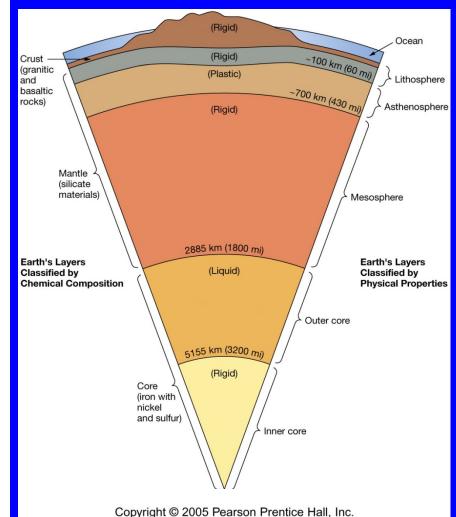


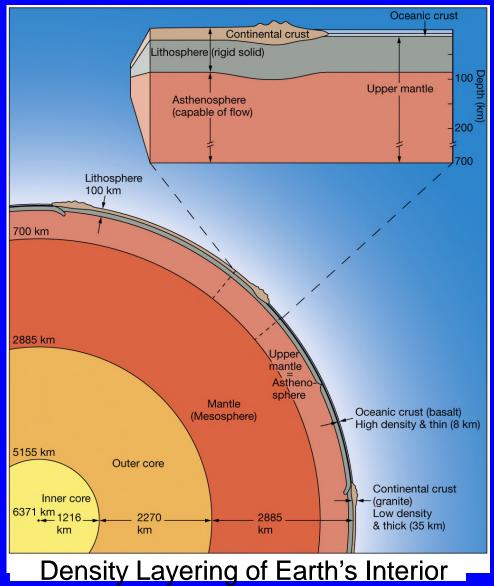
Early Earth likely entirely molten – gravitational segregation of dense metals (mostly Fe) to the center is the result.



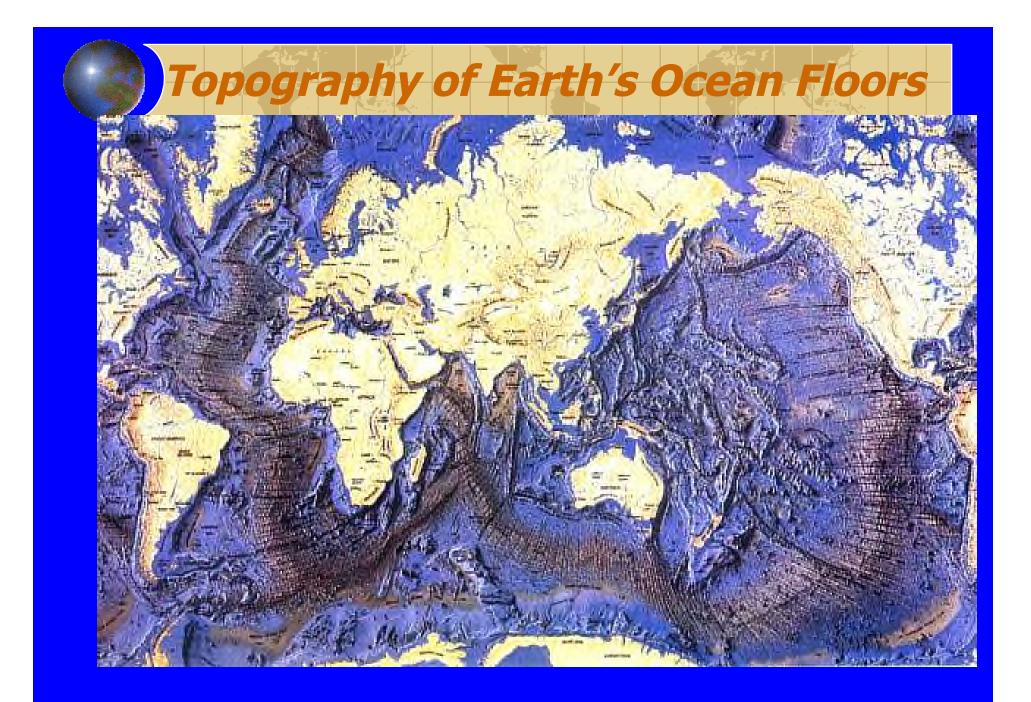
Earth's Anatomy Today

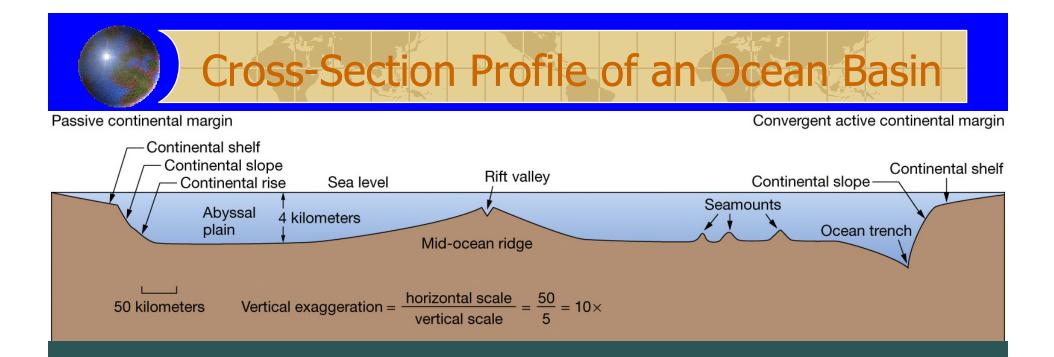
Chemical and Physical Nature of Earth's Interior







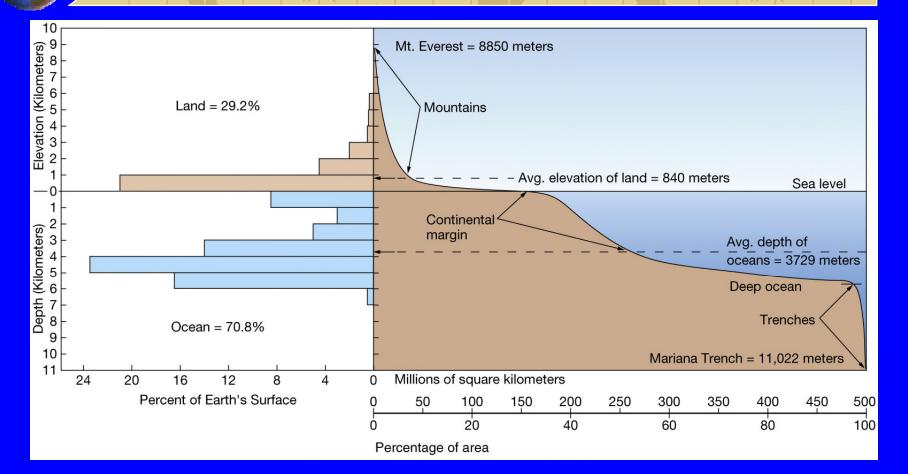




Large-Scale Ocean Bottom Features

- ✓ Continental shelf, slope, and rise
- ✓ Abyssal plains and hills
- ✓ Mid-ocean ridge and rift valley
- Oceanic islands, seamounts, and guyots
- ✓ Ocean trench

Elevation Relief Profile of Earth's Crust



- 1. Sea level
- **2.** Continental shelf
- **3. Continental slope**
- 4. The deep ocean floor
- 5. Mean depth of ocean 3700m
- 6. Mean altitude of land 840m
- 7. Mt. Everest 8848m
- 8. Mariana Trench 11022m

Earth's Continents and Ocean Basins

1) Two Different Types of Crust

- ✓ Continental Granitic
- ✓ Oceanic Gabbroic

2) Continental Crust

- ✓ Lighter (2.7 g/ml)
- ✓ Thicker (30 km)
- ✓ High Standing (1 km elev.)

3) Oceanic Crust

- ✓ Denser (2.9 g/ml)
- ✓ Thinner (7 km)
- ✓ Low Standing (- 4 km elev.)



1) Two Different Types of Crust

- \checkmark Continental = Granitic
- ✓ Oceanic = Gabbroic

2) Continental Crust

- ✓ Lighter (2.7 g/ml)
- ✓ Thicker (30 km)
- ✓ High Standing (1 km elev.)

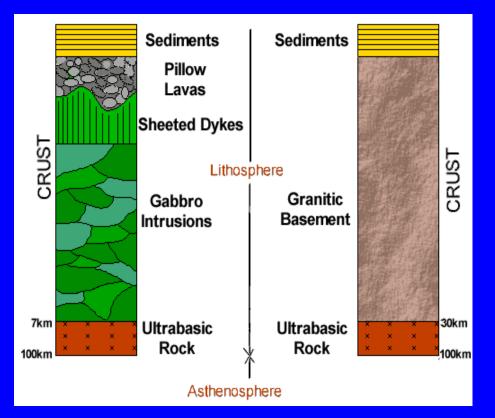
3) Oceanic Crust

- ✓ Denser (2.9 g/ml)
- ✓ Thinner (7 km)
- ✓ Low Standing (- 4 km elev.)

Oceanic Crust Gabbroic Rock

Two Primary Types of Earth Crust

Continental Crust Granitic Rock

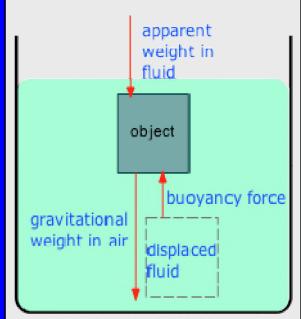


1) Buoyancy is an important force on objects immersed in a fluid.

2) Buoyancy is the fluid pressure exerted on an immersed object equal to the weight of fluid being displaced by the object.

Concept of Buoyanc

- 3) The concept is also known as Archimedes's principle
 - Principle applies to objects in the air and on, or in, the water.
 - Principle also applies to the crust "floating" on the mantle, which is specially termed "isostacy".
- 4) Density is a controlling factor in the effects of buoyancy between an object and its surrounding immersing fluid
 - The greater the difference in density between the object and the fluid, the greater the buoyancy force = sits high
 - The lesser the difference in density between the object and the fluid, the lesser the buoyancy force = sits low



Example of Buoyancy: Boat on a Lake

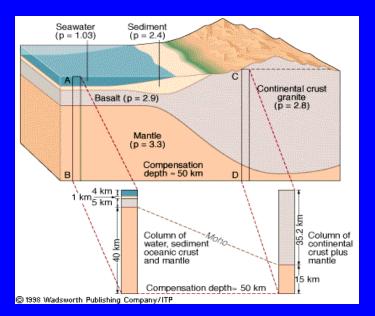


What is the density of the boat with cat in relation to the lake water?

The Concept of Isostasy

Defined: state of gravitational equilibrium between the earth's *rigid* lithosphere and *fluid* asthenosphere, such that the tectonic plates "float" in and on the underlying mantle at height and depth positions controlled by plate thickness and density.

The term "isostasy" is from Greek "iso" = equal; "stasis" = equal standing.



Earth's strong rigid plates exert a downward-directed load on the mobile, underlying weaker, plastic-like asthenosphere – pushing down into the mantle.

> The asthenosphere exerts an upward pressure on the overlying plate equal to the weight of the displaced mantle – *isostatic equilibrium* is established.

Mantle will flow laterally to accommodate changing crustal loads over time – this is called *isostatic adjustment*

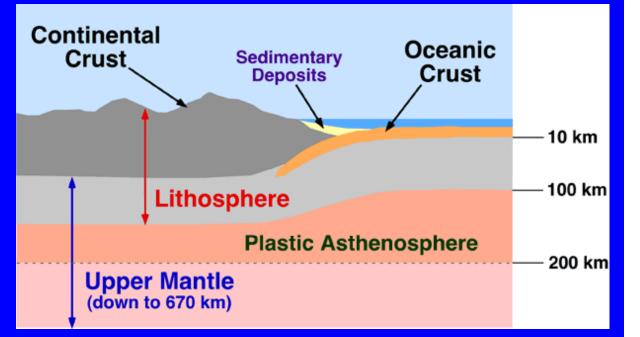
Plate tectonics, erosion and changing ice cap cause isostatic disequilibrium

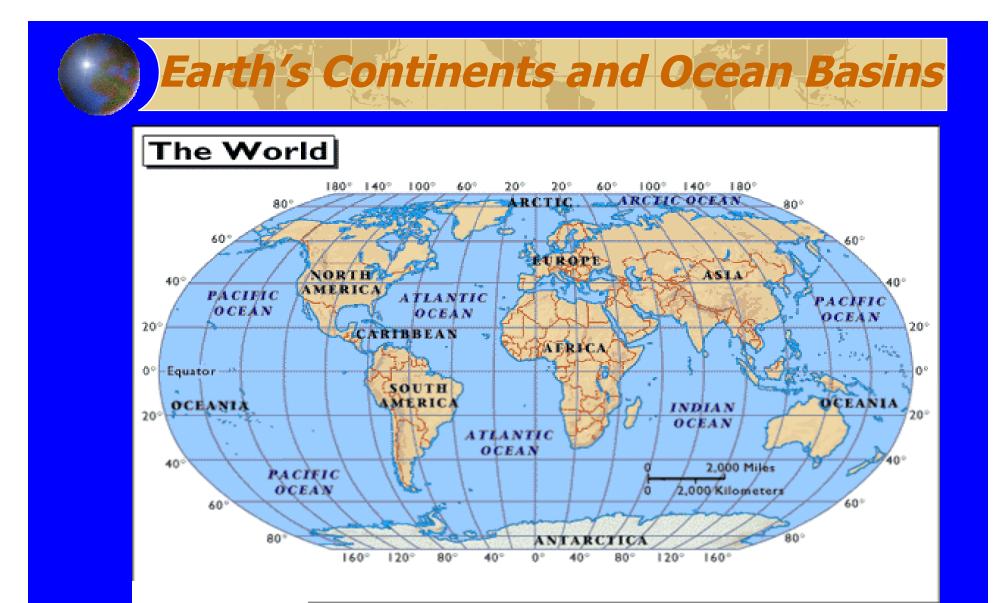
Crust Floating in Mantle

1) Isostatic Equilibrium Between Crust and Mantle; Lithosphere and Asthenosphere

2) Isostatic Adjustments Made Over Geologic Time When A Layer's Density and/ or Thickness Changes

4) Isostatic Adjustments Produce Vertical Movement of Crust – Uplift or Subsidence





Key Points: 1) Up until 50 years ago, all physiology maps of earth showed ocean basins as blue = lack of sea bottom data.
2) Continental land masses were well-mapped much earlier on.

OCEANOGRAPHY COMES OF AGE

- Technologic Innovations Light Up the Ocean Bottoms
 - ✓ Sonar and Radar Mapping
 - ✓ Piston coring and Drilling
 - ✓ Magnetometer surveys
 - ✓ Radiometric and fossil dating
 - ✓ Submersible investigations
 - ✓ Subsurface seismic surveys
 - ✓ Computer-assisted research

Detailed Seafloor Image Emerges

 Ridges, fracture zones, trenches

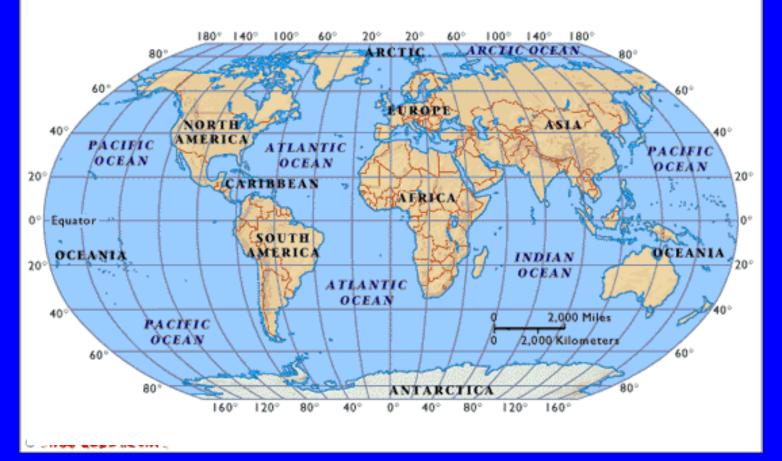
 Radical New Ideas Take Hold

 Seafloor Spreading and Subduction
 The Plate Tectonic Theory

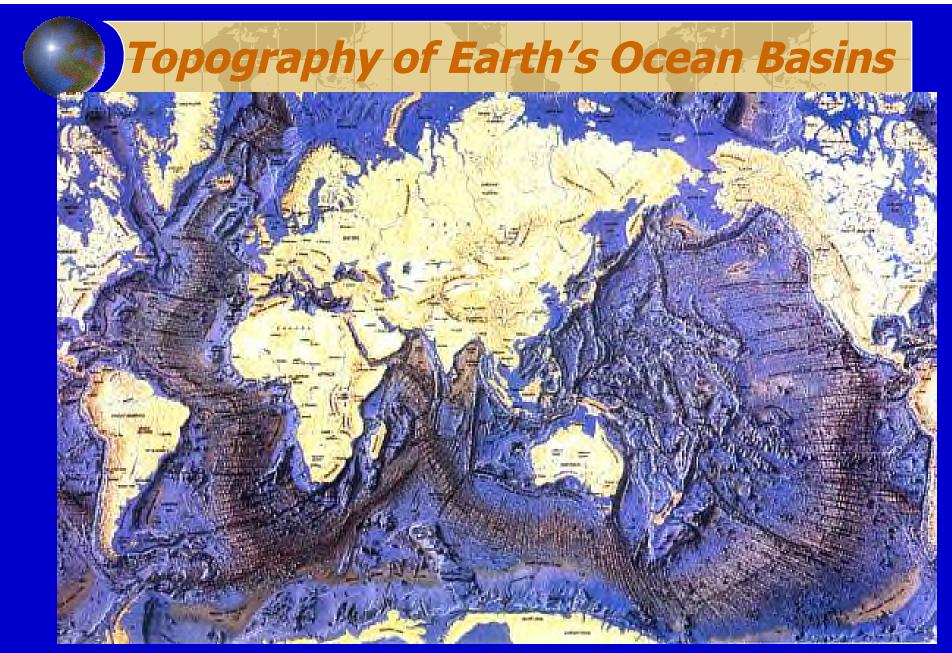


The Seafloor Illuminated!

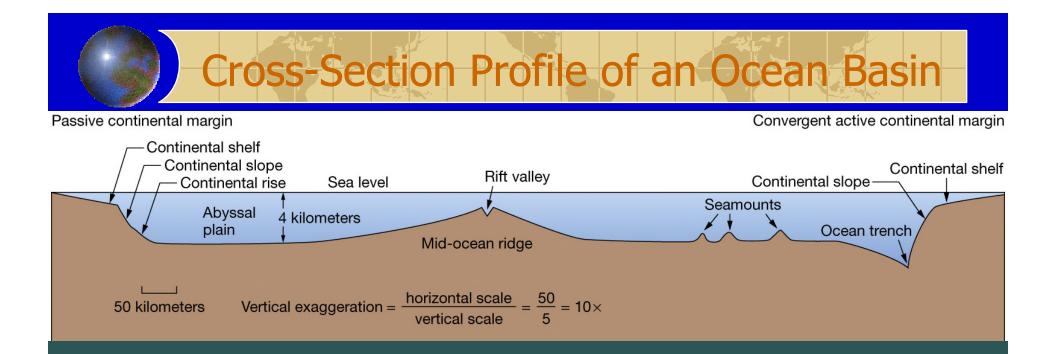
Earth's Continents and Ocean Basins



Typical Old-school World Map: Ocean Basins Colored Blue 100 Years ago: Unknown What Lied Under the Blue

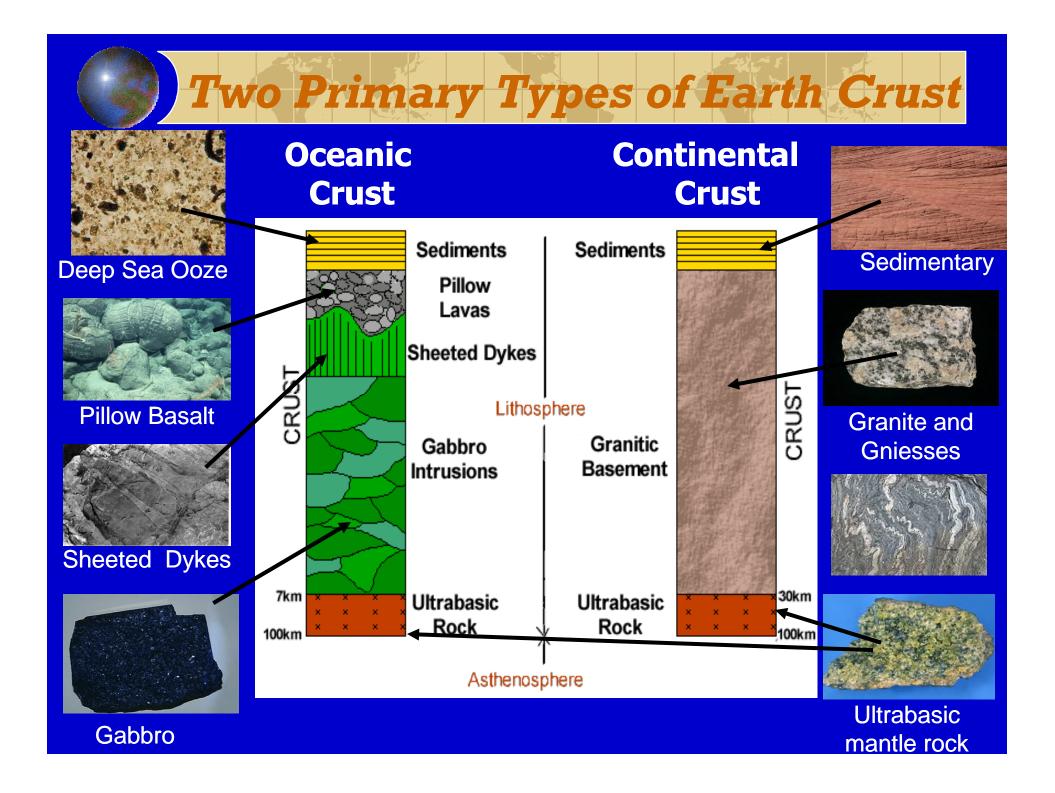


An Earth with No Ocean!



Large-Scale Ocean Bottom Features

- Continental shelf, slope, and rise
- ✓ Abyssal plains and hills
- ✓ Mid-ocean ridge and rift valley
- Oceanic islands, seamounts, and guyots
- ✓ Ocean trench



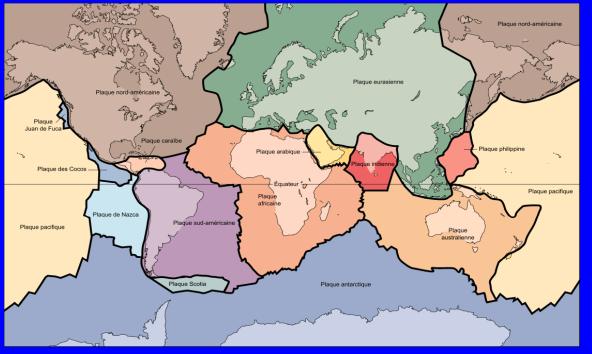
Earth's Layered Interior Oceanic crust **Chemical and Physical** Continental crust Lithosphere (rigid solid) Nature of Earth's Interior 100 🛛 Upper mantle Asthenosphere (km (Rigid) (capable of flow) Ocean 200 ~100 km (60 mi) (Rigid) Crust (granitic (Plastic) 00 and Lithosphere ~700 km (430 mi basaltic Lithosphere rocks) 100 km Asthenosphere (Rigid) 700 km Mantle (silicate materials) Mesosphere 2885 km (1800 mi) 2885 km Earth's Lavers Earth's Layers Uppei (Liquid) Classified by Classified by **Chemical Composition Physical Properties** mantle Astheno sphere Outer core Mantle Oceanic crust (basalt) (Mesosphere) High density & thin (8 km) 5155 km (3200 mi 5155 km (Rigid) Outer core Core (iron with Continental crust nickel Inner core and sulfur) (granite) 6371 km 1216. Low density Inner core 2270 2885 & thick (35 km) km km km **Density Layering of Earth's Interior** Copyright © 2005 Pearson Prentice Hall, Inc.

INTRO to PLATE TECTONIC THEORY

Key Features:

- ✓ 14 Lithosphere Plates
- 6 Major, 8 Minor
- ✓ 100-300 km thick
- Strong and rigid
- ✓ Plates float on partially molten asthenosphere
- ✓ Plates are mobile
- Cm's/yr motion rates
- ✓ Seafloor Spreading creates new oceanic plates

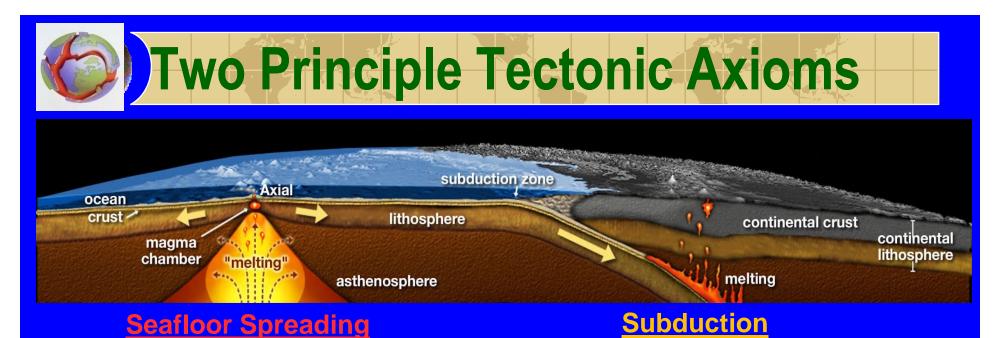
✓ Subduction destroys older oceanic plates



Earth's Lithospheric Plates

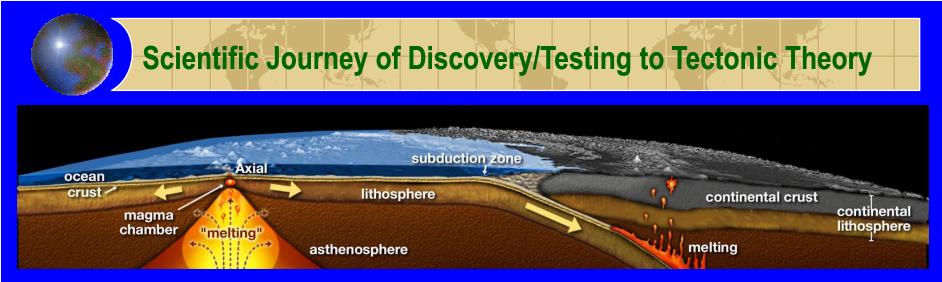


Animation of Overview of Plate Tectonics - on YouTube



- 1) Seafloor Spreading = Plate Constructive
 - Coincides with mid-ocean ridges
 - Divergent plate boundary
 - Tholiietic basaltic volcanism
- 2) Subduction = Plate Destructive
 - Coincides with deep sea trenches and volcanic arcs
 - Convergent plate boundary
 - Explosive Andesitic volcanism

Animation of Overview of Plate Tectonics – on YouTube



- 1) What sorts of observations were made and where? Data collected?
- 2) What sorts of technologies were developed and used?
- 3) How were hypotheses tested? Validated hypotheses turned into supporting evidence? Predictions made?
- 4) How were various established lines of evidence/ideas integrated to form the plate tectonic theory?
- 5) Road of discovery starts with the continental drift hypothesis starting back in early 1900's

Continental Drift Hypothesis

TETHYS

LAÚURASIA

TRIASSIC

200 million years ago

CRETACEOUS

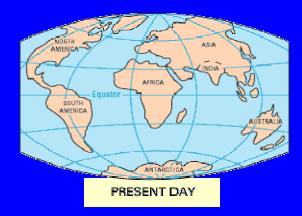
65 million years ago



PERMIAN 225 million years ago



JURASSIC 135 million years ago





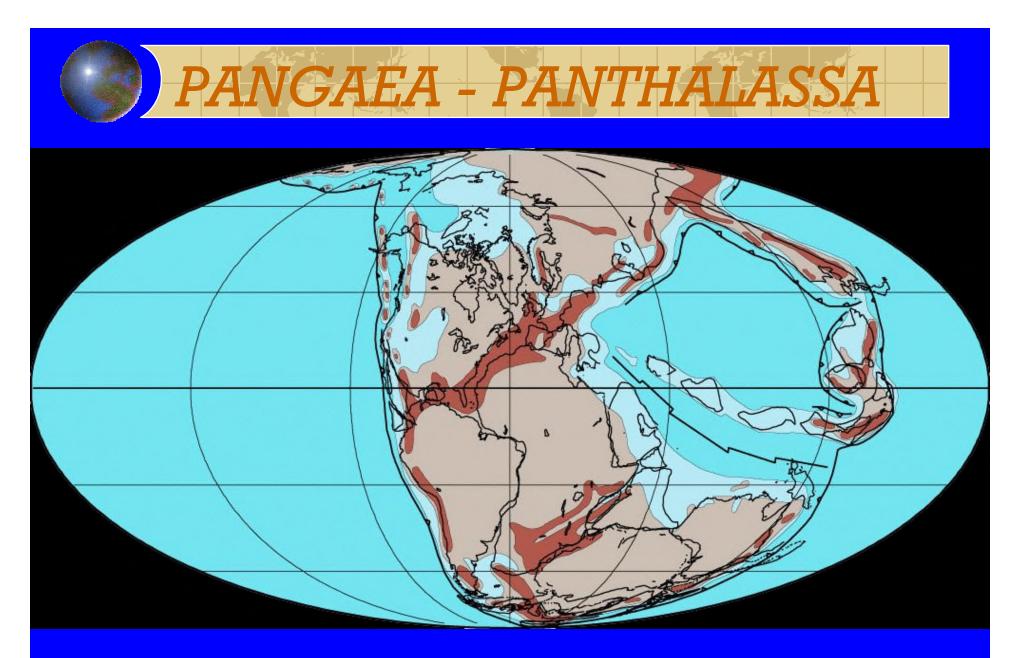
Alfred Wegener (1880-1930)

Main Ideas:

1. Alfred Wegener was the primary sponsor of hypothesis

Continental Drift Hypothesis

- 2. Supercontinent "Pangea" existed in the Permian Period
- 3. Pangea began to break up in the Triassic Period with dispersal, i.e. "drifting", of the rifted continents
- 4. Continental masses plowed through ocean crust
- 5. Strong lines of land-based evidence support the hypothesis
- 6. Driving mechanism for "continental drift" invalidated
- 7. Plate tectonics theory replaced continental drift idea



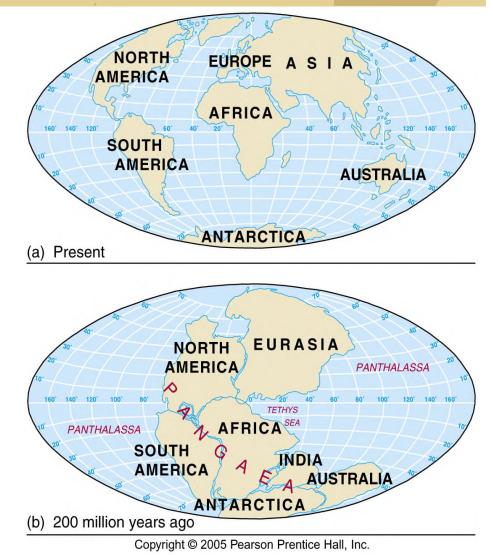
Permian Period - 220 Million Years Ago

Pangaea, Panthalassa, Triassic Breakup, and Continental Drift

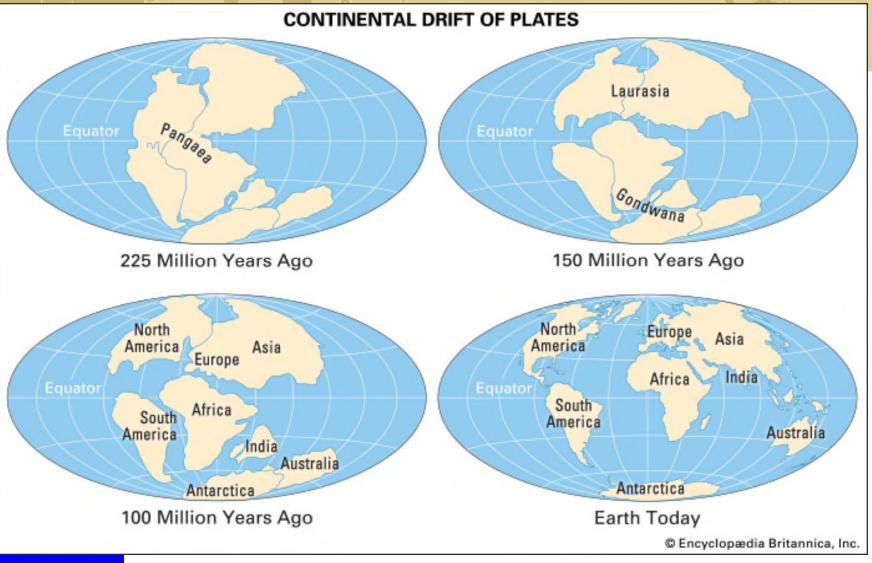
Animation shows the sequential breakup of the Pangea Supercontinent

The progressive breakup of Pangea occurred over the last 200 million years and will continue into the future

 Opening of Atlantic Ocean basin, collapse of Panthalassa Super-ocean basin, and Continental Drift

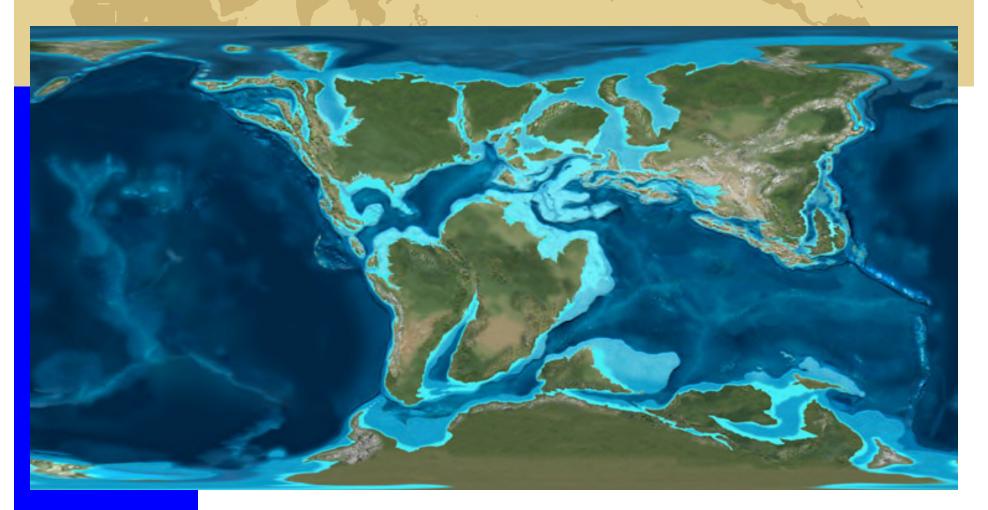


Pangaea and Continental Drift



Click Here for Animation of Breakup of Pangaea

Global Geology Through Time



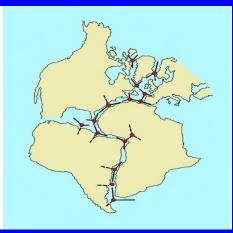
Click Here for Animation of Changing Global Geography Through Time

The Continental Drift Hypothesis

Wegener's Lines of Supporting Evidence:

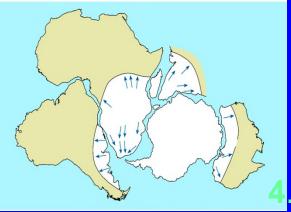
1. Fit of adjoining continental coastlines

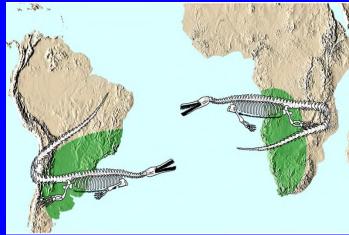




2. Truncated mountain and mineral belts

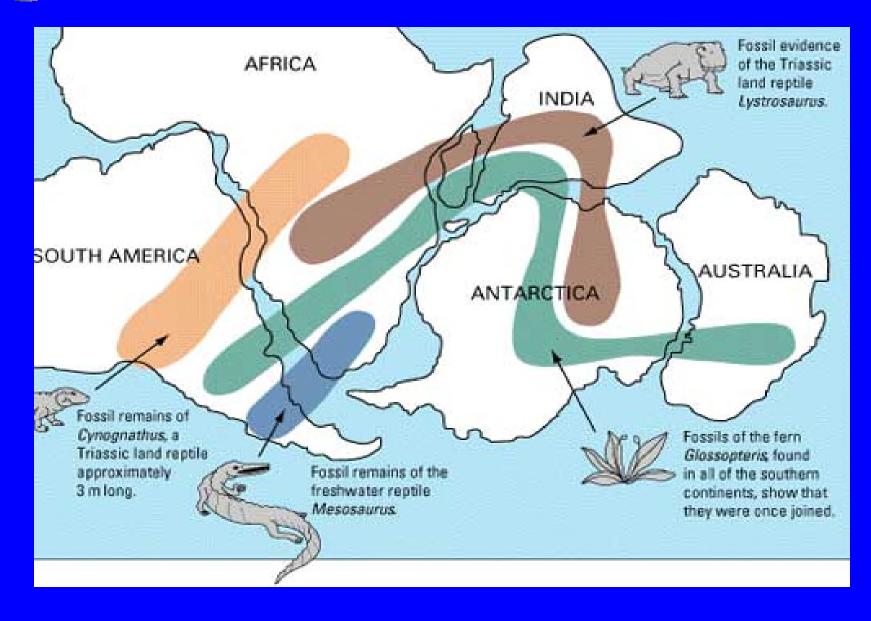
3. Intercontinental fossil affinities





Connection of ancient climatic belts

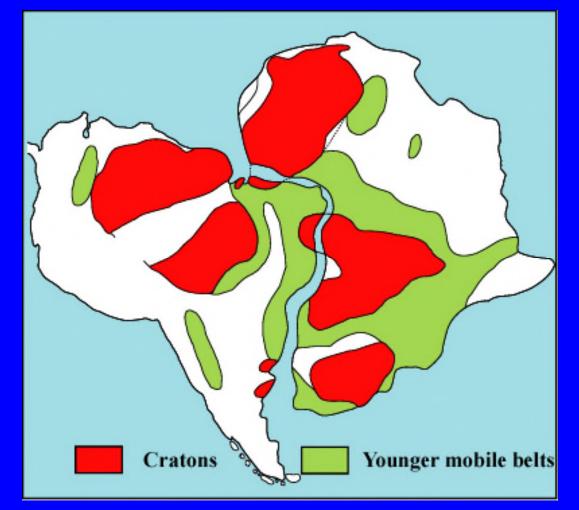
Gonwanaland Fossil Evidence



Gonwanaland Rock Evidence

Perfect Fit of Truncated:

- 1) Mountain Belts
- 2) Mineral belts
- 3) Terranes



Continental Drift Hypothesis



Conclusions

- Good land-based evidence for drift
- No evidence from ocean basins
- Driving mechanism invalidated
- No alternate drift mechanism found
- Hypothesis invalidated and nearly forgotten....until....????.



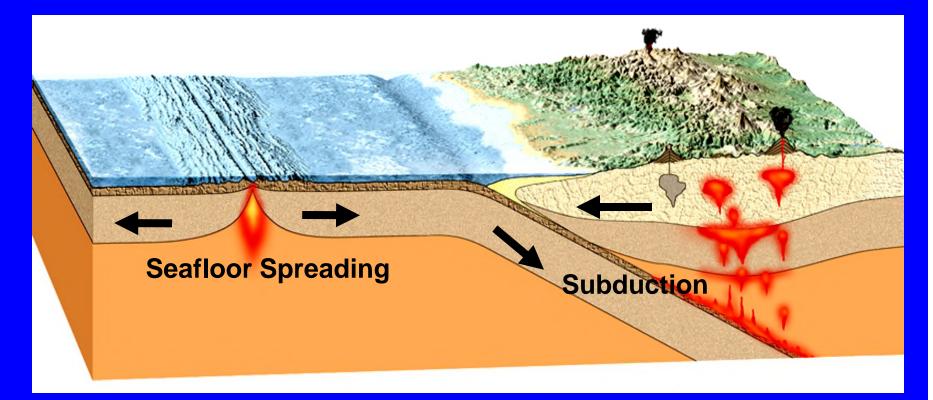
Wegener

Breakup of Pangea And Continental Drift

Two Principle Tectonic Processes

1) Seafloor Spreading = Plate Constructive

2) Subduction = Plate Destructive



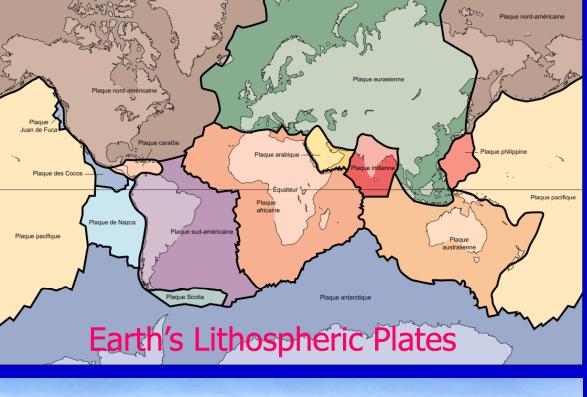
Animation of Overview of Plate Tectonics - on YouTube

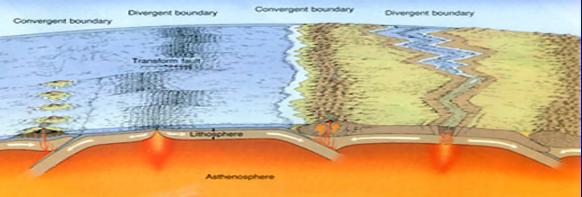
PLATE TECTONIC THEORY

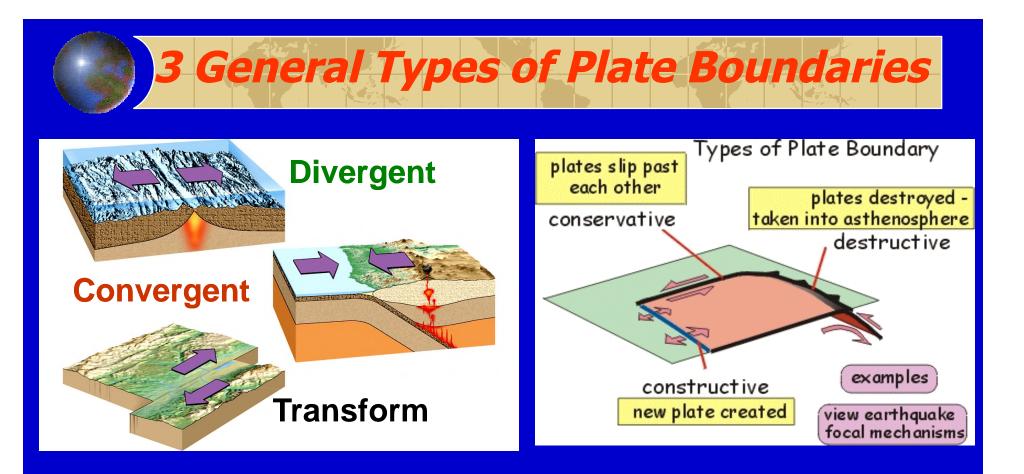
Key Features:

- ✓ 14 Lithosphere Plates
- 6 Major, 8 Minor
- ✓ 100-300 km thick
- Strong and rigid
- ✓ Plates float on partially molten asthenosphere
- ✓ Plates are mobile
- ✓ Cm's/yr motion rates
- ✓ Seafloor Spreading creates new oceanic plates

✓ Subduction destroys older oceanic plates





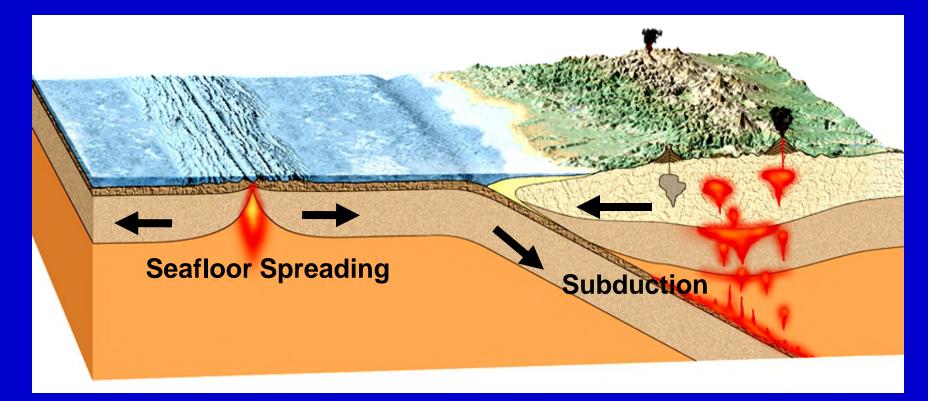


- 1) **Divergent = Constructive: creation of new oceanic plate**
- 2) Convergent = Destructive: destruction of old oceanic plate
- 3) Transform = Conservative: no creation or destruction of plates

Two Principle Tectonic Processes

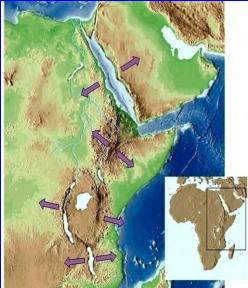
1) Seafloor Spreading = Plate Constructive

2) Subduction = Plate Destructive



Animation of Overview of Plate Tectonics - on YouTube

Divergent Plate Boundaries and Seafloor Spreading

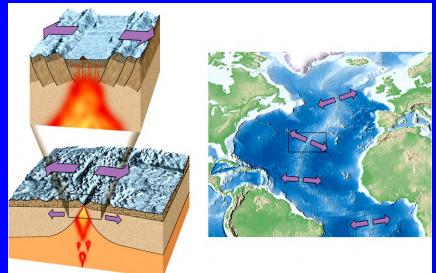


Main Ideas:

1. Seafloor spreading is a double conveyor beltlike process that produces "mirrored" growth of new seafloor between two diverging plates

2. Initiated by continental rifting event

3. Mid-ocean ridges are the most typical geographic expression of active spreading



 Plates "spread" apart to accommodate new additions at the ridge center (rift valley)

5. Basaltic magmas generated by the decompression melting of upwelling asthenosphere rock beneath the spreading centers

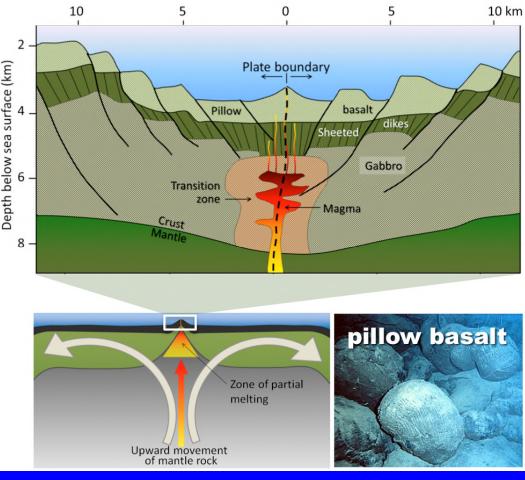
Key Features:

The illustration to the right shows the progressive growth of oceanic seafloor at a midocean ridge due to seafloor spreading

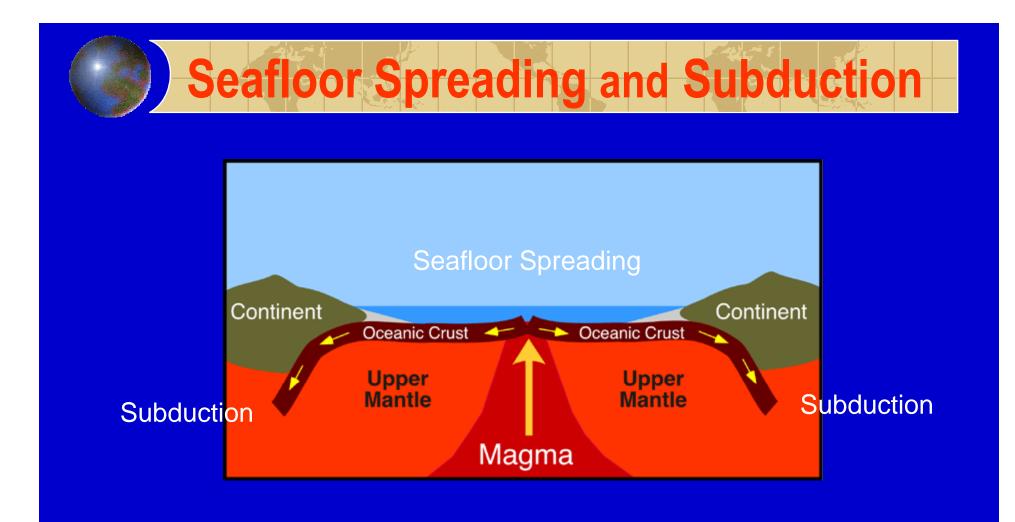
 Basaltic magmas arise from decompression melting of hot ascending asthenosphere beneath the mid ocean ridge

As new oceanic lithosphere is constructed at the mid ocean ridge, older plate material passively moves off and away from both sides of ridge

 Most oceanic lithosphere will eventually get subducted back into the asthenosphere Click Here for Seafloor Spreading Animation



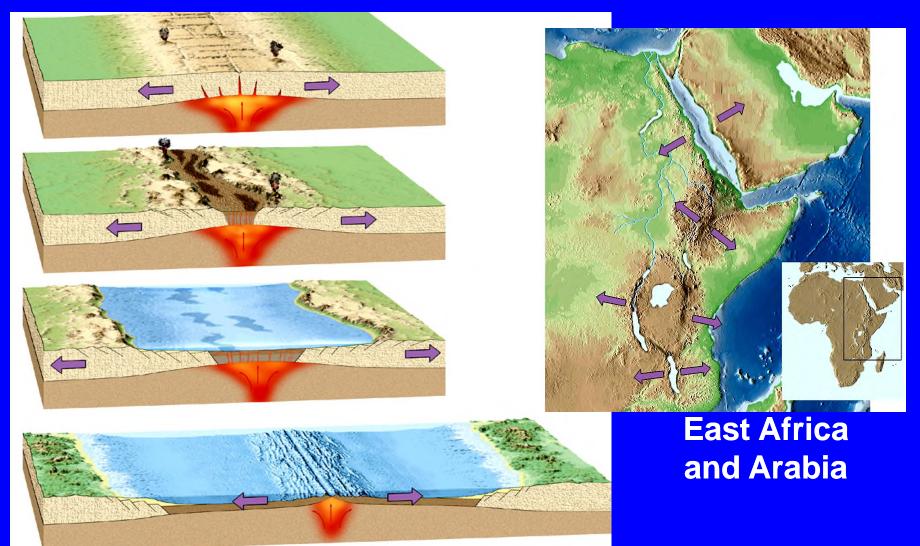
Seafloor Spreading Process

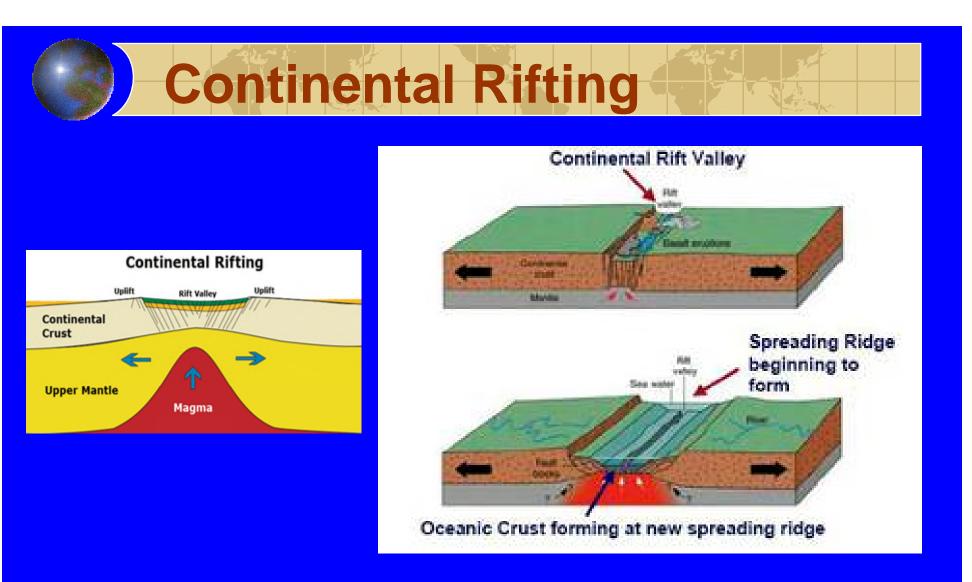


<u>Click Here for Animation of</u> <u>Seafloor Spreading and</u> <u>Subduction</u>

Continental Rifting & Ocean Basin Development

Progression from Continental Rifting to Seafloor Spreading





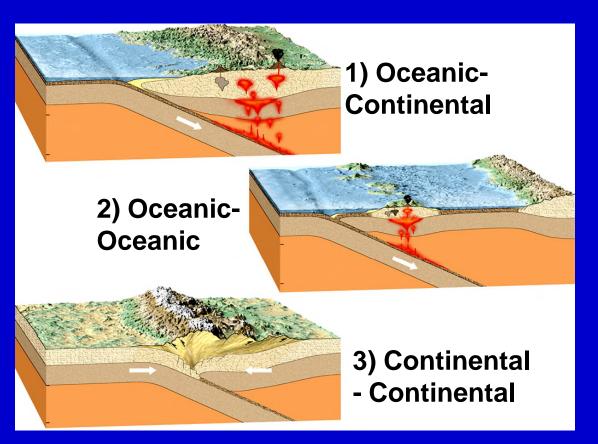
Click Here for Animation of Continental Rifting

3 Types of Convergent Plate Boundaries

1) Oceanic-Continental Subduction-related continental margin arc

2) Oceanic-Oceanic Subduction-related continental margin arc

3) Continental- Continental Collision boundary of two continents



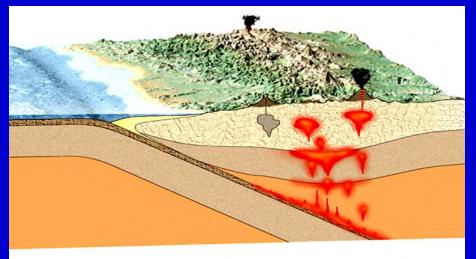
Key Points: Convergent plate boundaries are the sites of 1) formation of new continental crust, 2) intense crustal deformation and 3) recombination of continental masses.

Main Ideas:

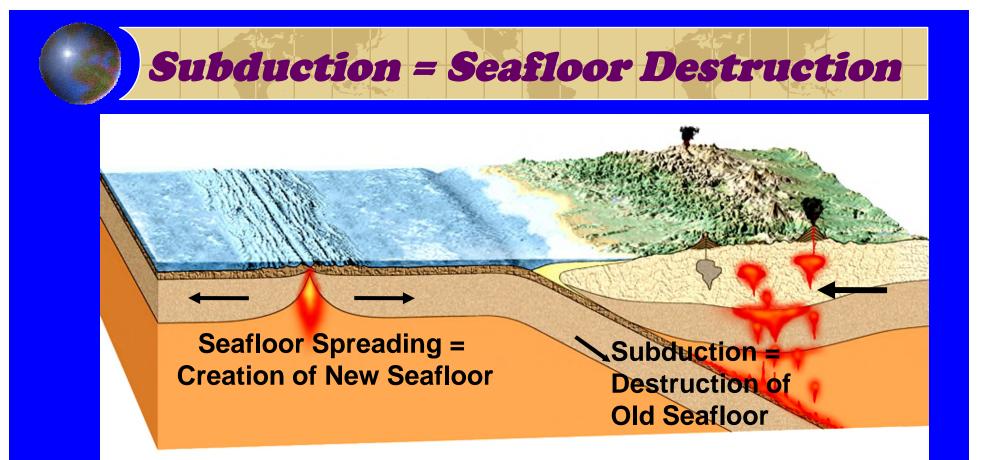
 Process of destroying old oceanic lithosphere by sinking down into the mantle at convergent plate boundaries

Subduction

2) Subduction zones are marked by a paired trench-volcanic arc system



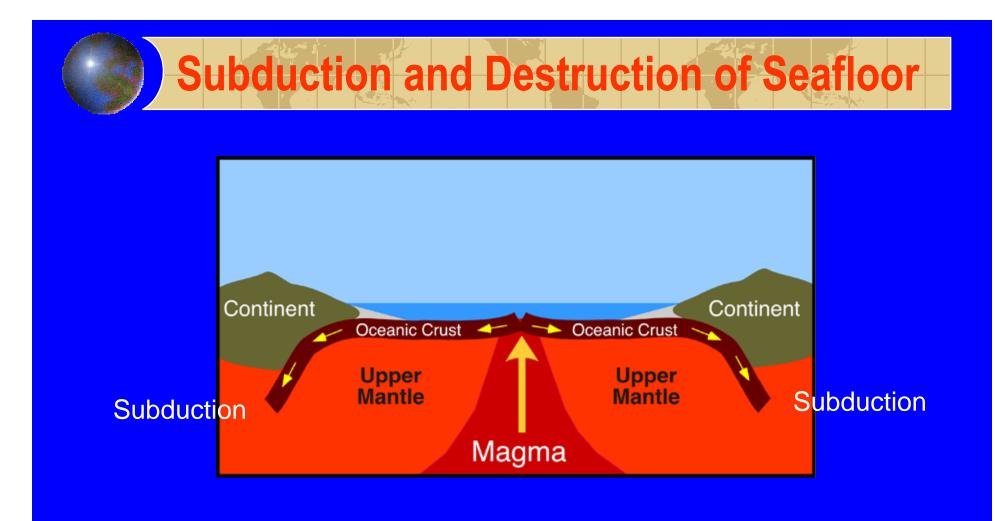
- **3)** Andesite-dominated volcanic arc H₂O magmas are generated by dehydration melting of subducted slab and mantle wedge beneath the volcanic arc
- 4) Highly explosive arc eruptions due to high silica, H₂O and CO₂ content
- 5) Subduction causes ocean basins to collapse
- 6) Subduction initiates the accretion of exotic, buoyant, crustal terranes
- 7) Subduction is the site where new continental crust is being created



1) Subduction is caused by over-dense oceanic plate sinking back into the asthenosphere under its weight = main driving force of plate tectonics.

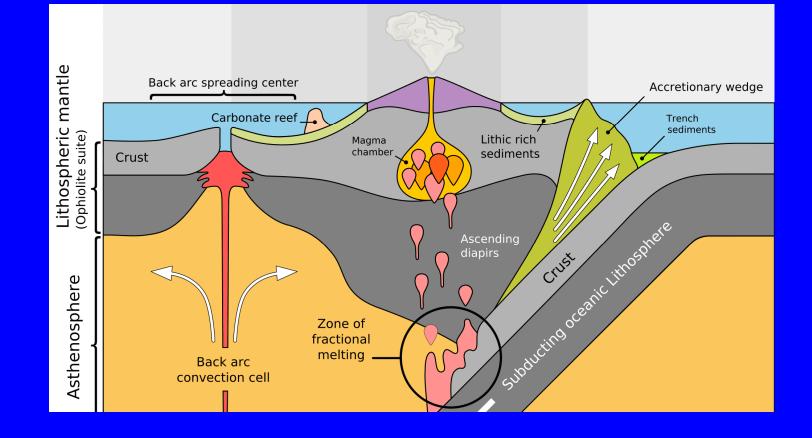
2) Interplate convergent motion at subduction zones leads to the diverging, pull-apart, seafloor spreading plate boundaries = ocean plate mass balance.

3) Seafloor spreading is the crustal mass counter-balancing process to the subduction of older density-unstable seafloor crust sinking back into mantle.



Click Here for Animation of Subduction

Formation of Volcanic Arcs



Subduction and Formation

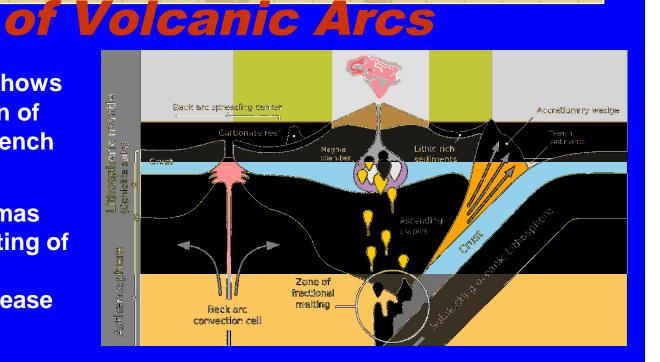
Key Features:

Illustration to the right shows the progressive destruction of old oceanic seafloor at a trench due to subduction.

 Water-rich basaltic magmas generated from partial melting of asthenosphere above the subducting slab, due to release of ocean water from slab

 Subduction-related magmas rise and intrude up through overlying plate creating a volcanic mountain chain or arc

 Other consequences of subduction are terrane accretion and collapsing ocean basins.



<u>Click Here for Animations of</u> <u>Volcanic Arc Formation</u>

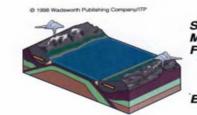
Subduction and Ocean Basin Collapse

Three Stages of Ocean Basin Collapse

- 1) **Declining =** Basin shrinkage
- 2) Terminal = MOR subducted

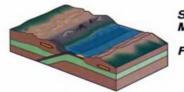
3) Suturing = Continental collision and extinguished subduction

The *climax* of an ocean basin collapse is the formation of a tall, extensive "fold and thrust" mountain chain, much like the Himalayas of today, along with the extinction of the subduction system (loss of active volcanism).

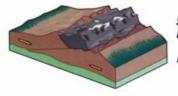


Stage: Declining Motion: Convergence Features: Subduction begins. Island arcs and trenches form around basin edge. Example: Pacific Ocean

17



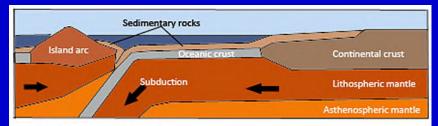
Stage: Terminal Motion: Convergence, collision and uplift Features: Oceanic ridge subducted. Narrow, irregular seas with young mountains. Example: Mediterranean Sea



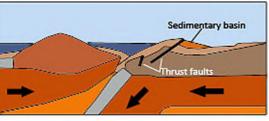
Stage:	Suturing
Motion:	Convergence and uplift
Features:	Mountains form as two continental crust masses collide, are compressed and override.
Example:	India-Eurasia collision. Himalaya mountains

Fig. 3-25 History of an ocean. (Second of two acetates.)

Volcanic Arcs and Terrane Accretion

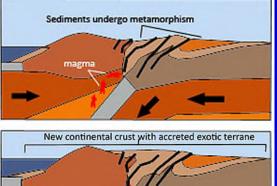


As subduction progresses, the island arc and continental crust resist being subducted and are pulled together, pinching sediments between them. Thrust faults develop.

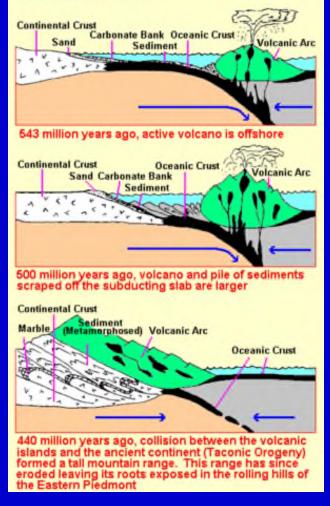


The oceanic crust subducts and the island arc collides with the continental crust, causing sediments to be metamorphosed and further development of thrust faults. Subducted oceanic crust melts in the mantle and plutons rise along the subduction pathway.

The island arc becomes sutured to the continental crust. Rising plutons become emplaced in the lower crust and drive additional metamorphism of sediments. Subduction below the suture zone ceases.



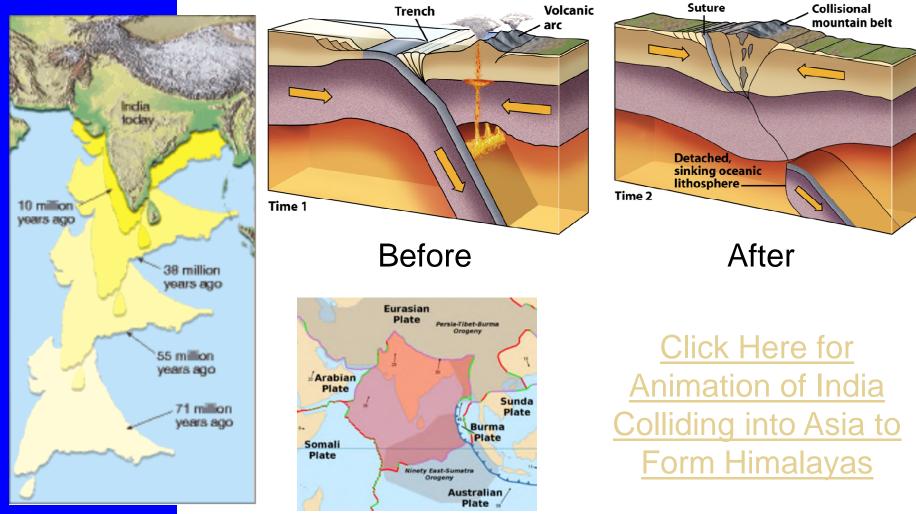
Cross Sections of Eastern North America (as it may have looked)



Click Here for Animation of Terrane Accretion

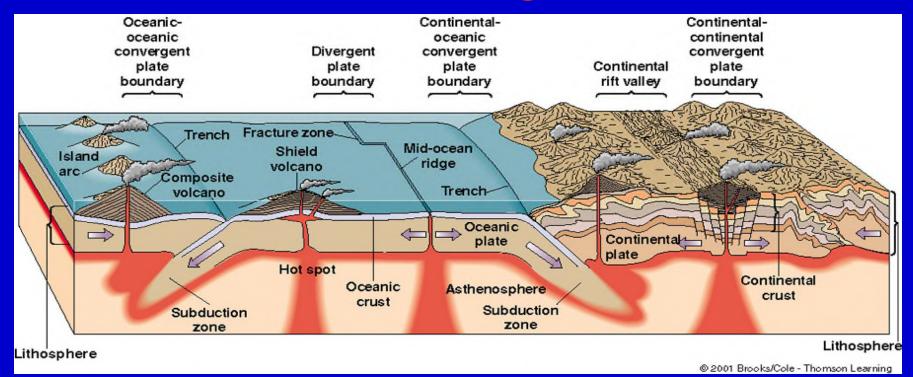


Continental Collision Events



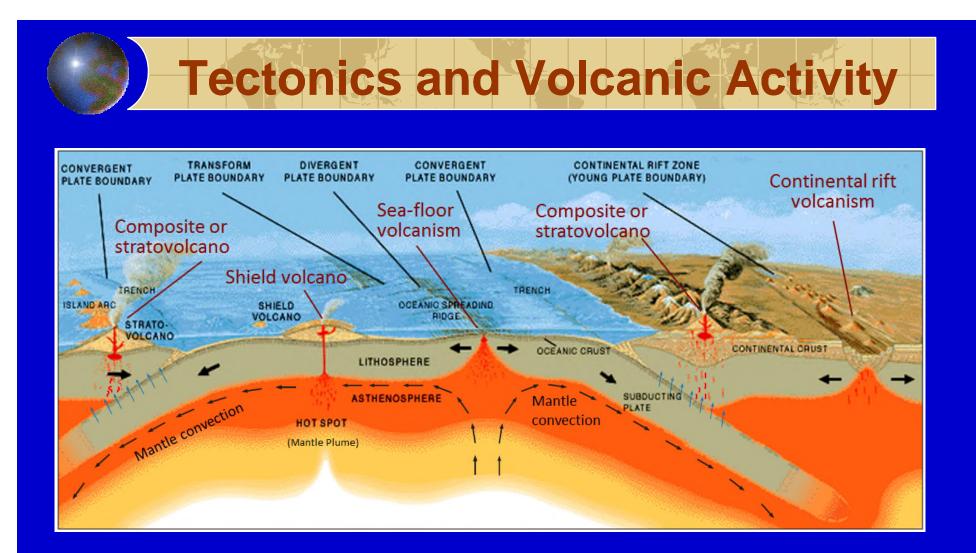
- Associated Faulting and Volcanism -

Plate Tectonic Boundaries

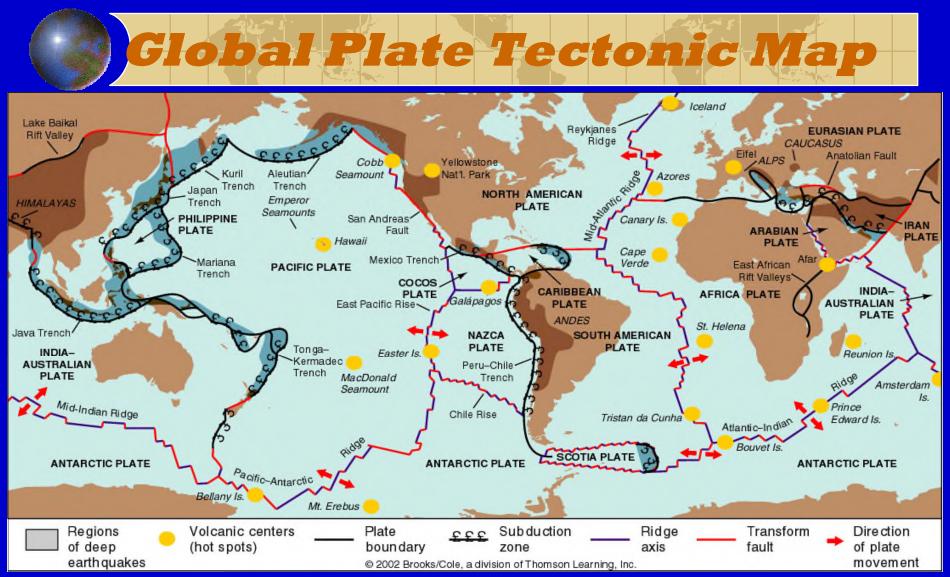


Questions:

How many types of plate boundaries do you recognize here?
 Which type of plate boundaries have little to no volcanism? Why?
 How does the plate tectonic theory explain inner-plate hot spots?



<u>Click Here for Animations of the Relationship Between</u> <u>Plate Tectonics, Volcanism, and Plate Boundaries</u>

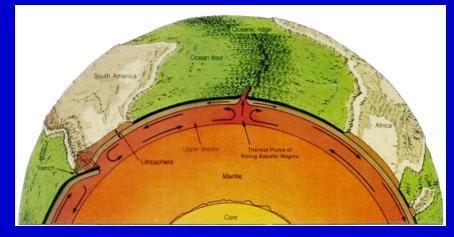


Key Points: 1) Each plate moves with a unique direction and speed 2) Fastest plates are those with greatest length of subducting edge. 3) Slowest plates have no subducting edges.



Subductive Thought/?





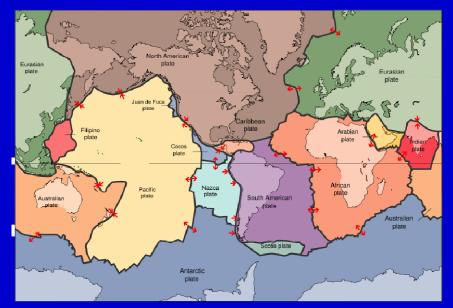
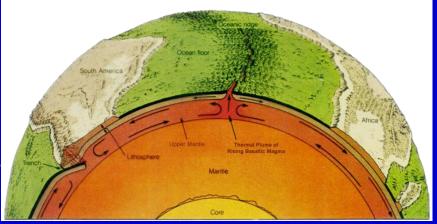


PLATE TECTONICS - Review

<u>Key concepts:</u>

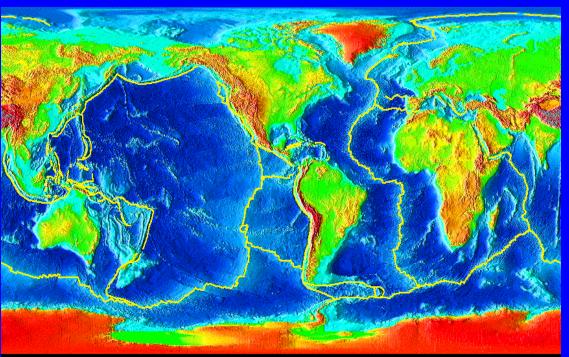
- 1) Earth's crust and uppermost mantle broken up into 18 mobile, rigid slabs called lithospheric plates
- 2) Lithospheric plates ride independently atop the underlying partially-molten mantle called the asthenosphere
- 3) Three types of dynamic lithospheric plate boundaries: Divergent, Convergent, and Transform
- 4) Divergent boundaries
 - Continental rifting
 - Seafloor-spreading
 - Creation of new oceanic plate
- **5)** Convergent boundaries
 - Subduction
 - Destruction of older oceanic pl
 - Terrane accretion
 - Continental collision
- **6)** Transform boundaries
 - Strike-slip faulting
- 7) Plate tectonics driven by density, heat and gravity (convection)
- 8) Plate tectonic theory explains most geologic phenomena



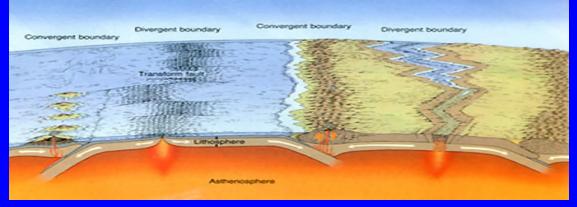
Review of Today's Topics

<u>Topics</u>

- ✓ Age of the Earth
- ✓ Earth Physiology
- ✓ Continental Drift
- ✓ Plate Tectonics Theory
- ✓ Seafloor Spreading
- ✓ Subduction
- ✓ Terrain Accretion



Crustal Plate Boundaries



Preparation for Next Meeting

Next Meeting: Plate Tectonics II

- Seafloor Spreading and Subduction
 Evidence for Plate Tectonics
 Plate Dynamics
- 4) Driving Mechanisms

Homework Assignment:

 Read Chapter 2, 13, 14 in Textbook
 Study the Instructor's Website
 www.geoscirocks.com
 Lecture Notes
 PowerPoint
 ER Videos 3, 4, 5, 6

